



## Full Length Article

## Seismicity characteristics of the Gulf of Aqaba seismogenic zone and their hazard implications in northwestern Saudi Arabia

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## ABSTRACT

The seismogenic characteristics of the Gulf of Aqaba zone have been assessed using the maximum likelihood method to estimate various earthquake recurrence parameters. These parameters encompass the  $\beta$ -value, annual recurrence rate ( $\lambda$ ), and maximum probable magnitude ( $M_{max}$ ). This assessment has identified three sub-seismogenic zones, each corresponding to specific structural faults within the Gulf. These zones are associated with the Aragonese, Arnona and Aqaba faults, delineating pull-apart basin structures in the Gulf of Aqaba. An updated earthquake catalogue has been compiled using a unified moment magnitude ( $M_w$ ) scale to improve the analysis, established by developing two empirical relationships. According to the findings of this study, there is a possibility that the Aragonese seismogenic zone could experience an earthquake with a maximum magnitude of 7.7, highlighting a significant seismic hazard in the region. While acknowledging the inherent uncertainties in this assessment, a probabilistic seismic hazard was calculated for hard rock conditions within a spatial area divided into elementary cells, each measuring  $0.1^\circ \times 0.1^\circ$ . The highest peak ground acceleration (PGA) is associated with a spectral frequency of 5.0 to 10.0 Hz and could significantly impact building codes in the region. The spatial distribution variations of seismic hazard corresponding to the proposed sub-seismogenic zones indicate a high degree of crustal heterogeneity and seismotectonic complexity. This comprehensive assessment contributes to understanding seismic hazards that may import from the Gulf of Aqaba seismogenic zone.

## 1. Introduction

Saudi Arabia's vision for 2030 includes significant development projects in the northwest region, such as constructing an international bridge connecting Saudi Arabia and Egypt. Given the scale of critical infrastructure development in the region, understanding, estimating, and mitigating earthquake hazards is paramount.

The extent of damage caused by earthquakes to structures depends on various factors, including peak ground acceleration (PGA), ground motion intensity, structural design and foundation, soil rigidity, liquefaction potential, site geology, and the structure's location relative to the earthquake epicenter (Sengezer et al., 2008). Seismic hazard assessments are vital for accurately estimating potential property and life losses in risk analyses. These assessments determine site-specific expected ground motions near critical infrastructure like railways,

highways, bridges, and dams. A crucial aspect of a comprehensive seismic hazard analysis is defining seismogenic zones and characterizing seismic hazards, which involves understanding the spatial distribution of peak ground accelerations (PGAs), response spectra, exceedance probabilities, and return periods.

From the perspective of Saudi Arabia, the Gulf of Aqaba seismogenic zone is a significant concern as it is considered the most earthquake-prone area in the Red Sea region. The region experienced two significant earthquakes in August 1993 and November 1995. The 1993 earthquake with a local magnitude of 5.8 (Abdel Fattah et al., 1997) occurred near the western coastline of the Gulf along the Aragonese/Arnona fault and likely contributed to the 1995 earthquake with a moment magnitude of 7.2 (Hofstetter et al., 2003). The 1995 earthquake caused extensive damage in northwest Saudi Arabia, followed by intense aftershocks. This seismic event significantly impacted and destroyed the

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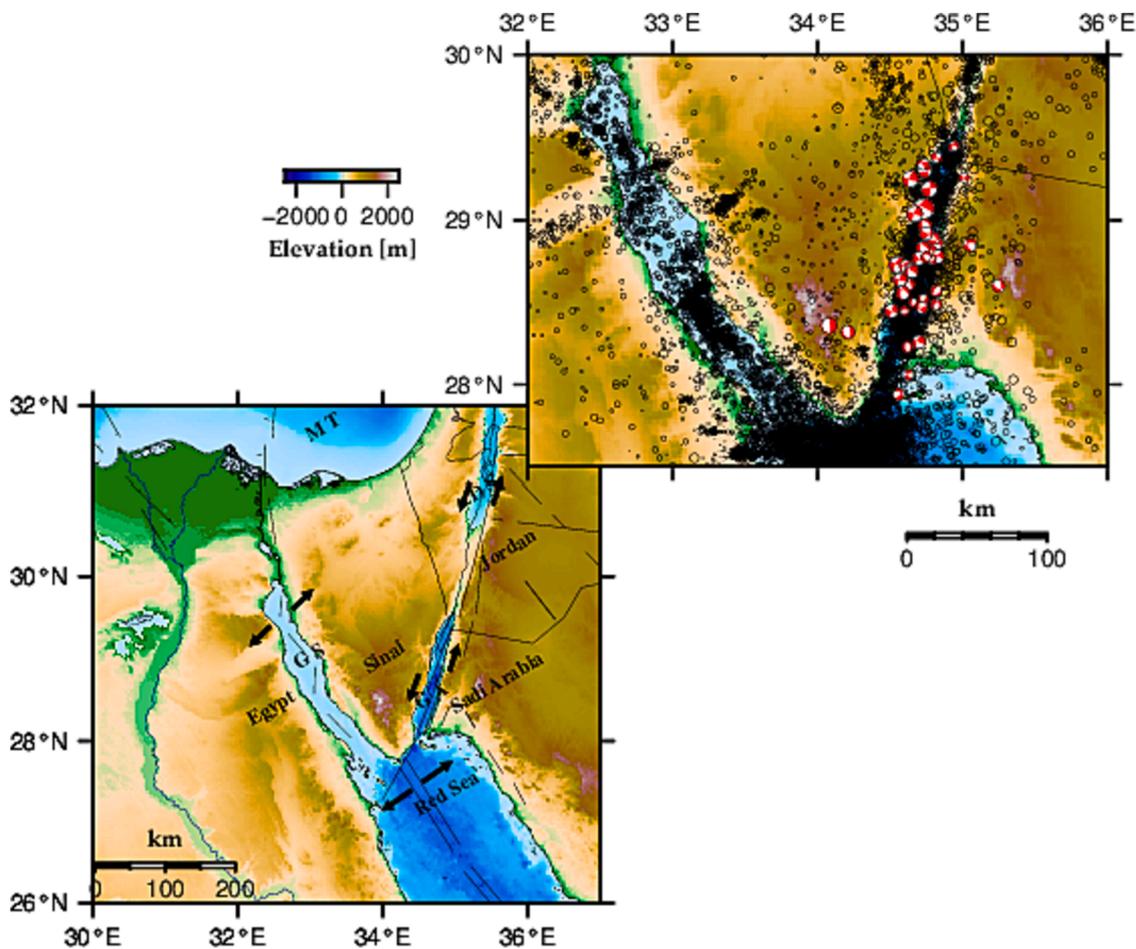


Fig. 1. Tectonic settings and seismicity distribution in the Gulf of Aqaba and its vicinity region.

coastlines of nations adjacent to the Gulf. Severe destruction was witnessed in Nuweiba, Egypt, the city nearest the epicenter, with most buildings suffering extensive or complete destruction (Klinger et al., 1999). It is essential to highlight that the Gulf of Aqaba earthquake source zone poses a potential seismic threat to neighboring regions, including Saudi Arabia, Egypt, Jordan, and Palestine. The rapid development of megacities near this seismogenic zone, like Neom in Saudi Arabia, may raise concerns about potential hazards.

Seismogenic models describing the spatial and temporal distribution of potential seismic sources are crucial to seismic hazard assessment. In the Gulf of Aqaba, the tectonics are influenced by three major strike-slip faults: the Aqaba Fault, the Aragonese Fault, and the Arnona Fault, which generate pull-apart basin structures. The presence of these sub-seismogenic zones was not thoroughly considered in previous assessments but has been incorporated into the seismic source modeling in this study. It is important to note that earlier studies have indicated that the highest PGAs in the Gulf of Aqaba seismogenic zone tend to align parallel to the coast and decrease in magnitude as one moves away from the coast.

This study generated hazard spectra on a hard rock site in Neom, Saudi Arabia, for various return periods based on the characteristics of earthquakes in the Gulf of Aqaba seismogenic zone. Earthquake data from 1964 to 2023 was collected and analyzed to update probabilistic seismic hazard maps on both sides of the Gulf. The methodology used is based on established techniques for estimating fundamental earthquake hazard parameters, including maximum regional magnitude ( $M_{max}$ ), magnitude-frequency relationship ( $\beta$  value), and mean seismic activity rate ( $\lambda$ ). These parameters help enhance our understanding of seismic hazards, aid earthquake preparedness and mitigation and are calculated

using the maximum likelihood estimation method. We used the method introduced by Kijko and Sellevoll (1989, 1992) and Kijko (2004) to estimate the maximum regional magnitude  $M_{max}$  and other related earthquake hazard parameters.

## 2. Tectonics and seismicity

The Gulf of Aqaba seismogenic zone is situated near the transform fault that connects the Red Sea rifting system with the Dead Sea transform fault system, as previously reported by Mechie and El-Isa (1988). However, the current tectonic dynamics in this seismogenic zone are primarily influenced by the ongoing separation of the Arabian and African plates. This separation significantly affects the rotational movement of the Sinai sub-plate, as Salamon et al. (2003) discussed. Fig. 1 visually represents the ongoing tectonic processes and the spatial distribution of epicenters in the northern Red Sea region from 750BCE to 2020 CE. Additionally, Fig. 1 presents the spatial distribution of focal mechanism solutions for the earthquakes listed in Table 1. The arrangement of hypocenters in terms of latitude and longitude indicates that the Gulf of Aqaba seismogenic zone spans approximately 120 km in length and 15 km in width. Most of the seismic events within this zone occur at depths ranging from 5 to 20 km. According to the International Seismological Centre, the mainshock of the 1995 earthquake sequence, with a moment magnitude ( $M_w$ ) of 7.2, is the most significant known earthquake in the region. However, the Gulf of Aqaba seismogenic zone experiences relatively infrequent earthquakes with magnitudes exceeding 6.5, as noted by Klinger et al. (1999).

Seismic activity in the Gulf of Aqaba seismogenic zone is characterized by isolated seismic sequences occurring spatially and temporally,

Table 1

List of focal parameters for earthquakes shown in Fig. 1, as taken from Almadani (2020).

ID	Date	To	Lat [°]	Long [°]	d [km]	M	Strike [°]	Dip [°]	Rake [°]	Mo	Mw
1	19820323	104800	27.90	34.30	10	4.7M <sub>L</sub>	220	65	-040	4.42E+22	4.4
2	19830203	134604	29.19	34.77	24	5.3M <sub>w</sub>	360	80	014	2.31E+23	5.3
3	19851231	194241	28.85	35.05	09	4.8M <sub>L</sub>	155	60	-030	6.03E+22	4.5
4	19890909		28.57	34.82	10	4.1M <sub>L</sub>	205	50	-110	6.84E+21	3.8
5	19930703	233410	28.86	34.82	18	4.7M <sub>L</sub>	114	88.69	149	4.42E+22	4.4
6	19930803	4305	28.73	34.55	17	6.0M <sub>w</sub>	007	62	-117	2.63E+24	6.0
7	19930803	124305	28.73	34.55	17	6.0M <sub>L</sub>	294	87.27	-157	2.51E+24	5.6
8	19930803	163323	28.36	34.08	15	5.7M <sub>L</sub>	356	79.41	-083	9.89E+23	5.7
9	19930807	45540	28.61	34.63	10	4.2M <sub>L</sub>	217	50.19	-042	9.33E+21	3.9
10	19930820	230959	28.72	34.61	02	4.6M <sub>L</sub>	116	79.84	140	3.24E+22	4.3
11	19931103	183932	28.7	34.65	07	4.9M <sub>L</sub>	086	76.13	-148	8.22E+22	4.5
12	19931108	10602	28.69	34.65	08	4.7M <sub>L</sub>	303	80.49	-120	4.42E+22	4.4
13	19931204	233411	28.89	34.8	10	4.6M <sub>L</sub>	093	68.07	-148	3.24E+22	4.3
14	19951122	1526	29.07	34.73	18	7.2M <sub>w</sub>	196	59	-015	1.71E+26	7.2
15	19951122	124704	29.3	34.74	15	5.0M <sub>L</sub>	074	80.04	-150	1.12E+23	4.6
16	19951122	221657	28.32	34.21	15	5.2M <sub>L</sub>	357	60.54	-069	2.09E+23	4.8
17	19951123	717	29.25	34.64	10	5.7M <sub>w</sub>	199	77	007	9.27E+23	5.7
18	19951124	164345	28.97	34.74	10	4.9M <sub>L</sub>	165	54.28	-068	8.22E+22	4.5
19	19951211	13208	28.92	34.75	19	5.0M <sub>L</sub>	302	84	147	1.12E+23	4.6
20	19960103	100526	28.6	35.25	10	4.8M <sub>L</sub>	113	59.77	-138	6.03E+22	4.5
21	19960108	1318	29.38	34.82	06	3.8M <sub>L</sub>	340	78.86	164	2.69E+21	3.6
22	19960116	61700	29.34	34.73	06	4.3M <sub>L</sub>	159	82.79	149	1.27E+22	4.0
23	19960204	72300	29.45	34.94	06	3.6M <sub>L</sub>	294	77.43	-149	1.45E+21	3.4
24	19960221	5951	28.8	34.78	10	5.3M <sub>w</sub>	132	30	-104	2.31E+23	5.3
25	19970510	147	28.26	34.7	10	4.6M <sub>w</sub>	114	89	150	2.02E+22	4.6
26	20000308	142226	28.83	34.73	15	4.9M <sub>w</sub>	182	48	-048	5.74E+22	4.9
27	20000308	122229	28.64	34.57	10	4.8M <sub>w</sub>	198	31	-018	4.06E+22	4.8
28	20000308	142225	28.77	34.7	07	4.9M <sub>L</sub>	135	85.1	-169	8.22E+22	4.5
29	20000406	63734	28.78	34.83	12	4.8M <sub>L</sub>	0209	85.1	-012	6.03E+22	4.5
30	20000803	142225	28.64	34.57	10	4.8M <sub>L</sub>	108	82.31	167	6.03E+22	4.5
31	20010207	33900	29.26	35.01	21	3.6M <sub>w</sub>	044	80	-005	6.25E+20	3.6
32	20021110	5945	28.23	34.62	16	3.9M <sub>L</sub>	180	47.3	-084	3.67E+21	3.6
33	20040922	120023	28.45	34.6	10	3.2M <sub>L</sub>	270	64.05	-076	4.17E+20	3.0
34	20050120	134100	28.5	34.66	19	3.2M <sub>L</sub>	041	43	-056	4.17E+20	3.0
35	20080302	170600	28.76	34.79	14	3.2M <sub>L</sub>	058	75	-029	4.17E+20	3.0
36	20080404	140500	28.78	34.75	07	3. M <sub>L</sub>	145	46	-067	1.97E+21	3.5
37	20111021	123700	28.52	34.73	09	3.7M <sub>L</sub>	148	49	-147	1.97E+21	3.5
38	20130610	84400	28.06	34.63	06	3.9M <sub>L</sub>	142	73	-015	3.67E+21	3.6
39	20150627	53404	29.04	34.67	22	5.6M <sub>w</sub>	200	82	002	6.55E+23	5.6
40	20160516	14200	28.47	34.72	07	4.3M <sub>L</sub>	157	52	-173	1.27E+22	4.0
41	20160516	4559	28.45	34.51	18	5.2M <sub>L</sub>	019	88	-003	1.63E+23	4.7
42	20160813	30600	28.49	34.82	19	3.9M <sub>L</sub>	332	47	-153	3.67E+21	3.6
43	20161129	113	28.56	34.59	18	4.7M <sub>w</sub>	188	59	-044	2.86E+22	4.7
44	20170519	141600	27.94	34.56	04	4.1M <sub>L</sub>	304	50	-027	6.84E+21	3.8

as described by Abdel-Fattah et al. (2006). Notable earthquake sequences have occurred in various years, including 1983, 1990, 1993, 1995, 2015, and 2016, as documented by Almadani (2020). The distribution of fault plane solutions across the Gulf of Aqaba has revealed different earthquake faulting mechanisms. Normal earthquake faulting mechanisms are observed in the southernmost parts of the Gulf of Aqaba and the northern region of the Red Sea. In contrast, earthquake activity is primarily associated with strike-slip faulting mechanisms in the central and northernmost regions of the Gulf of Aqaba, highlighting the significant role of strike-slip faulting in this area. This transition in fault mechanism styles from south to north suggests that faulting processes in the northern Red Sea do not directly extend into the Gulf. Seismic tomography analysis conducted by El Khrepy (2016) did not uncover evidence of oceanic crust development within the Gulf.

### 3. Earthquake catalogue

An earthquake catalogue from 750BCE to 2020 CE was compiled for the study area. Data sources include historical earthquake records reported by Ambraseys et al. (2005) and instrumental earthquake data from various sources, including Maamoun et al. (1984), the International Seismological Center (ISC) data file, Harvard Seismology, the USGS/NEIC catalogue, ISC Bulletin, and catalogue, Preliminary Determination of Epicenters (PDE) from the National Earthquake Information

Centre (NEIC), Bulletins of the Egyptian National Seismic Network, and the Saudi Geological Survey (SGS). The distribution of earthquakes was assumed to follow a Poisson distribution, as Cornell (1968) described. In order to maintain the accuracy and reliability of the seismic activity rate, dependent events like foreshocks and aftershocks were excluded from the catalogue to avoid short-term anomalies. The algorithm proposed by Reasenberg (1985) was employed to remove the effects of non-background earthquakes.

Primary magnitude data in terms of body wave magnitude (Mb) and local magnitude (ML) were obtained from the ISC to calibrate moment magnitude (Mw). The ML values were calculated using the empirical relation by Shapira (1988), which was initially associated with sample ML values derived from short-period amplitude measurements on seismograms simulated to a Wood-Anderson instrument. Data for mL and mb were collected from the ISC and the National Earthquake Information Centre (NEIC), while corresponding Mw data were sourced from various providers, including CMT Harvard (Ekström et al., 2012), the Regional Centroid Moment Tensor catalogue (RCMT), and the Relative Moment Tensor catalogue (RMT).

In order to mitigate discrepancies resulting from the utilization of various instrument types, seismic wave measurements, and methods for calculating magnitude, earthquake magnitudes were standardized by their conversion to moment magnitude (Mw). Empirical relationships were developed to relate different magnitude scales (ML and mb) pro-

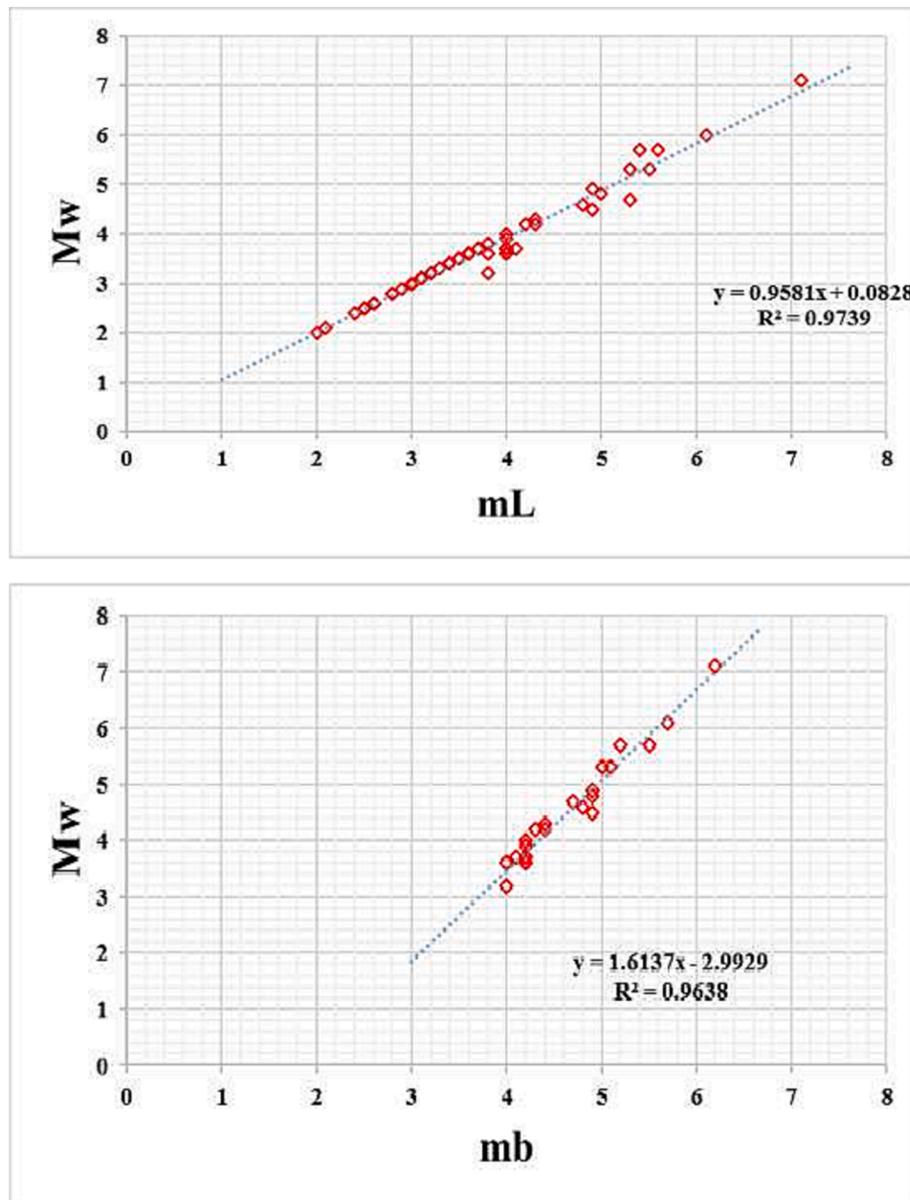


Fig. 2. Displayed the magnitude conversion relationships derived in this study.

vided by various agencies (Fig. 2a and b), and these empirical relationships are expressed as follows:

$$M_w = 0.96 \pm 0.02 \bullet ML + 0.083 \pm 0.08 \quad (1)$$

$$M_w = 1.61 \pm 0.07 \bullet mb - 2.99 \pm 0.31 \quad (2)$$

The method described by Stepp (1972), which relies on analyzing how the cumulative seismic activity's slope changes over time, was employed to determine the completeness magnitude levels of the earthquake catalogue (Mc). Following this technique, the earthquake catalogue was considered comprehensive for magnitudes above  $M_w$  3.0 since 1984,  $M_w$  4.5 since 1965, and  $M_w$  5.0 since 1900. Fig. 3 shows the spatial distribution of epicenters for the earthquake catalogue compiled in this study.

#### 4. Seismicity model

An essential aspect of assessing seismic hazards involves constructing seismogenic models that detail seismic sources' spatial and temporal distribution. These models encompass critical parameters, such as

seismicity rate, maximum earthquake magnitude, earthquake depth distribution, and the spatial arrangement of seismic sources. These parameters are estimated by leveraging diverse data sources, including historical earthquake records, geological mapping, and geophysical investigations. In the Gulf of Aqaba, the tectonics are influenced by three major strike-slip faults, each exhibiting a minor component of normal motion alongside their primary strike-slip movement. In this study, the seismic source modeling takes these faults into account.

A combination of extensional and shear deformations appears to govern the current tectonic processes in the Gulf, resulting in the development of a Neogene structure consisting of a series of pull-apart basins that make up the Gulf of Aqaba (Ben-Avraham et al., 1979; Tibor et al., 2010; Abdel-Fattah et al., 2016). These basins include the Eilat Deep, Aragonese Deep, Arnona Deep, and Dakar Deep, from north to south. These basins have varying thicknesses and are bounded by three major strike-slip faults: the Aqaba fault, the Aragonese fault, and the Arnona fault. The spatial extension of the Gulf of Aqaba Fault model and area seismic model is depicted in Fig. 4a and b.

According to Ribot et al. (2021), the Aragonese fault spans approximately 53 km, the Arnona fault covers around 83 km, and the Aqaba

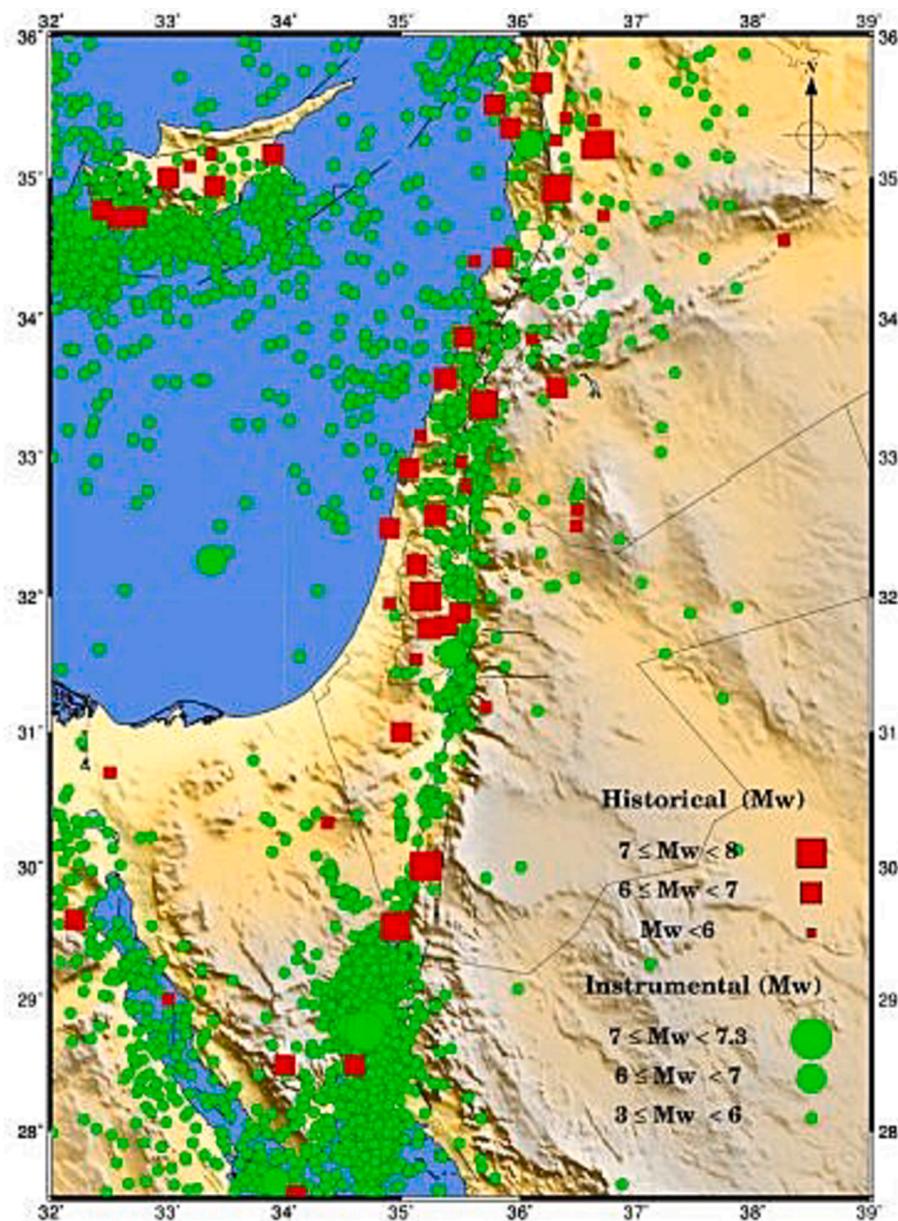


Fig. 3. The compiled earthquake catalogue for the Gulf of Aqaba and its surroundings.

fault extends over 160 km. The slip rates for these faults were determined by Bungum (2007) using GPS data from stations around the Gulf of Aqaba, revealing slip rates of 4.5 mm per year for the Aragonese fault, 4.5 mm per year for the Arnona fault, and 4.7 mm per year for the Aqaba fault. In order to determine the highest anticipated magnitude ( $M_{max}$ ) along each fault, recent scaling relationships from Thingbaijam et al. (2017) were employed. The seismicity parameters for these three faults are summarized in Table 2. These faults are incorporated into the first seismogenic model (Fig. 4a), which was developed to depict potential earthquake activity in the Gulf of Aqaba. The second model, the Aqaba Area seismic model (Fig. 4b), was created to account for seismicity associated with these faults using area-based seismogenic models. The seismicity parameters for these sub-seismogenic zones describe the likelihood of various magnitude earthquakes projected to occur in the region, with the distribution of earthquake magnitudes assumed to follow the Gutenberg and Richter (1954) relationship, as per Cornell (1968):

$$\log(n) = a - bm \tag{3}$$

where ‘n’ represents the cumulative number of earthquakes with a magnitude greater than or equal to ‘m,’ ‘a’ stands for the seismicity rate, and ‘b’ denotes the potential seismicity encompassing small and large earthquakes. The relationship expression is limited to the range between the lower threshold magnitude ( $M_{min}$ ) and the maximum expected magnitude ( $M_{max}$ ), with earthquakes more minor than  $M_{min}$  excluded from consideration. The mean annual rate of exceedance (represented as  $\lambda M$ ) is calculated using a doubly truncated exponential relationship,

$$\lambda M = \alpha \frac{\exp[-\beta(M - M_{min})] - \exp[-\beta(M_{max} - M_{min})]}{1 - \exp[-\beta(M_{max} - M_{min})]} \tag{4}$$

where  $\alpha = N(M_{min})$ , and  $\beta = b \cdot \ln 10$ .

The activity parameters ( $\beta$  and  $\lambda$ ) in the previous relationship were determined through the maximum likelihood estimation method developed by Weichert (1980), while  $M_{max}$  was derived using the statistical approach described by Kijko (2004). The seismicity parameters for the seismic models we proposed are provided in Table 3.

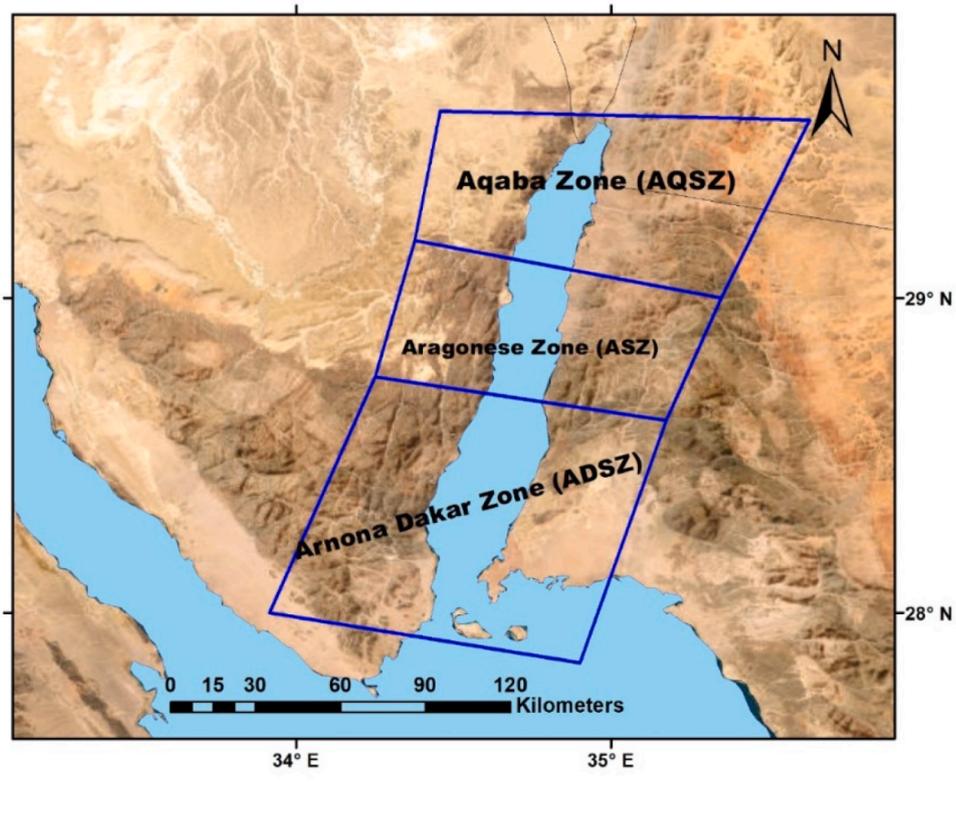
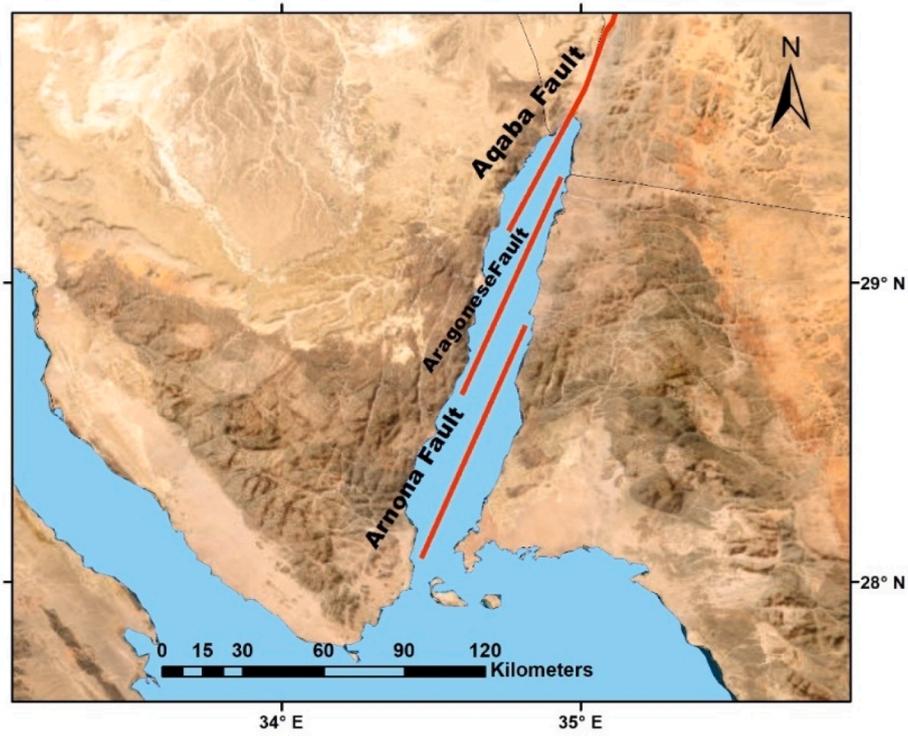


Fig. 4. Maps show the spatial extension of the Gulf of Aqaba (a) Fault model and (b) Area seismic model.

**Table 2**  
The estimated Mmax, slip rate, Mo, and a value at each fault segment.

Fault Segment	Length (Km)	Mmax	Slip rate (mm/y)	Mo (pa)	a value
Arnona	83	7.1	4.5	2.64E + 17	4.43
Aragonese	53	6.9	4.5	1.68E + 17	4.34
Aqaba	160	7.6	4.7	3.01E + 18	4.30

**5. Ground-motion models**

Ground motion prediction equations (GMPEs), commonly called ground motion models, were employed to estimate the magnitude of ground motion resulting from the seismic source characteristics. These models are developed based on data from strong-motion recordings of previous earthquakes. However, due to the limited availability of such records in Saudi Arabia, creating region-specific ground motion models presented a challenge. Consequently, ground motion models from areas with similar tectonic characteristics were used to predict the expected ground motion in the study area.

Following the recommendations provided by Cotton et al. (2006) and Boomer et al. (2010), four Ground Motion Prediction Equations (GMPEs) were selected to model ground motion in the current study, utilizing the logic tree approach. This approach considers epistemic uncertainty by accounting for variations in GMPEs. The chosen models include those proposed by Abrahamson et al. (2014), Akkar and Boomer (2010), Chiou-Youngs (2014), and Zhao et al. (2006).

**6. Probabilistic seismic hazard calculations**

The traditional approach to conducting a probabilistic seismic hazard analysis (PSHA) was initially introduced by Cornell (1968) and has subsequently been expanded upon and widely adopted by several researchers (e.g., Bender and Perkins, 1987). The likelihood of earthquake

events, as modeled by the Poisson distribution, can be mathematically expressed as follows:

$$P_n[m > M, t] = \frac{(\lambda t)^n \exp(-\lambda t)}{n!} \tag{5}$$

In the context of seismic hazard analysis, Pn [m, M, t] denotes the probability of experiencing ‘n’ earthquakes with magnitudes greater than ‘M’ within a specified area over a period ‘t’ while ‘λ’ represents the expected number of earthquake occurrences per unit of time in that area. The Poisson model assumes that earthquakes are spatially and temporally independent, that earthquake occurrences are consistent within a short time, and that two earthquakes will not happen simultaneously. The anticipated number of earthquake occurrences within a particular area can be estimated using statistical data from the site, based on the frequency-magnitude relationship initially defined by Gutenberg and Richter (1954). By taking the inverse logarithm of equation (4) and substituting it into the Poisson model, we get:

$$P_n[m > M, t] = \frac{\{\exp(\alpha - \beta M)\}^n \exp\{-[\alpha - \beta M]t\}}{n!} \tag{6}$$

From the above expression, we can find the probability of having no earthquake with a magnitude greater than M over a while t in each area as

$$P[m \leq M, t] = P_o[m > M, t] = \exp\{-[\alpha - \beta M]t\}. \tag{7}$$

The probability of having at least one earthquake with a magnitude greater than M in time t as

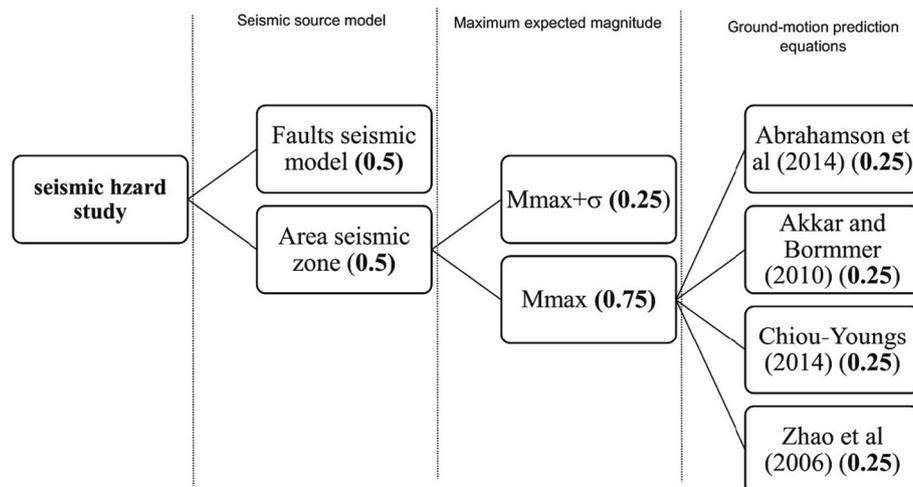
$$P[m > M, t] = 1 - P[m \leq M, t] = 1 - \exp\{-[\alpha - \beta M]t\} \tag{8}$$

The return period of an earthquake with a magnitude greater than M is

$$T_R = \frac{1}{\lambda} = \exp(\beta M - \alpha) \tag{9}$$

**Table 3**  
The seismicity parameters for each seismic source within the proposed seismic source model.

Zone No.:	Name of Zone	M <sub>min</sub> .	β	σ β	λ	M <sub>max</sub> <sup>obs</sup>	M <sub>max</sub>	σ M <sub>max</sub>
1	Arnona Dakar Zone (ADSZ)	3.5	2.1	0.16	0.49	6.1	6.7	0.20
2	Aragonese Zone (ASZ)	3.0	2.05	0.15	0.75	7.2	7.7	0.50
3	Aqaba Zone (AQSZ)	3.0	2.07	0.07	0.45	5.1	5.6	0.50



**Fig. 5.** The logic tree approach used in the current study. A bold number between brackets is the weight.

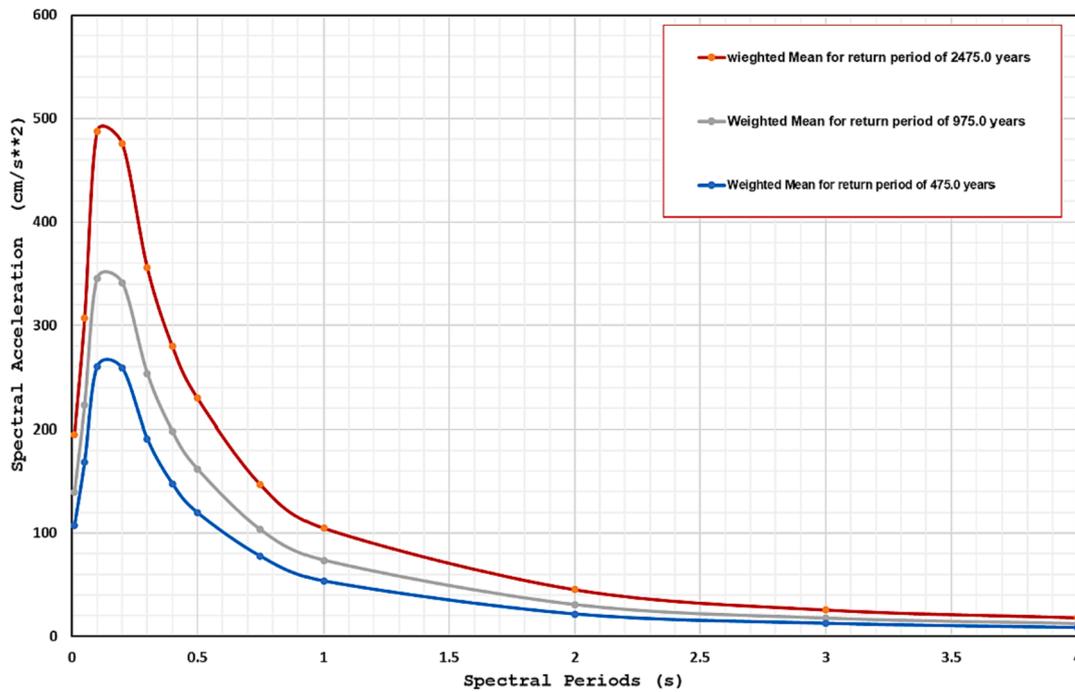


Fig. 6. Uniform hazard spectra at the rock site in Neom City for return periods of 475, 975 and 2475 years.

## 7. Logic tree

A logic tree approach addressed the inherent epistemic uncertainty in quantifying PSHA. This approach involved considering alternatives for seismic source models, standard deviations of model values describing maximum expected magnitudes for each seismic source, and ground motion attenuation relationships. The logic tree assigns weights to various branches to account for the likelihood of each of the three alternatives.

The logic tree comprises three components in the present analysis, as illustrated in Fig. 5. The first component deals with the uncertainty in seismic source models and has two branches. The first branch pertains to the Aqaba fault-source model, while the second is for the area seismic source model. The assessment of this component concludes that while it is plausible to evaluate hazards for individual faults, the resulting hazards are highly influenced by variations in fault seismic characteristics due to uncertainties. These uncertainties could lead to overestimating the hazard, particularly in the near-source region. Therefore, an equal weight is assigned to the fault-source and area seismic source models in the developed logic tree in this study. The second component addresses the uncertainty in estimating the maximum expected magnitude, where the expected mean maximum magnitude carries a weight of 0.75, while a weight of 0.25 accounts for deviations from the mean maximum magnitude ( $M_{max} + \sigma$ ). The third component deals with alternative equations for modeling the expected ground motion. Four Ground Motion Prediction Equations (GMPEs) are adapted and assigned equal weighting.

The widely accepted software code EZ-FRISK performs calculations according to the logic tree. Computations are implemented using a grid resolution of  $0.1^\circ \times 0.1^\circ$ , covering the northwestern part of Saudi Arabia and the Sinai Peninsula. The calculations generated seismic hazard curves, uniform hazard spectra (UHS), and hazard maps illustrating PGA. The current analysis considers various spectral periods (0.5, 0.1, 0.2, 0.3, 1.0, 2.0, 3.0, and 4.0 s) along with three levels of ground shaking (with probabilities of exceedance of 10 %, 5 %, and 2 % over 50 years). These spectral periods and different levels of ground shaking are commonly used in building codes.

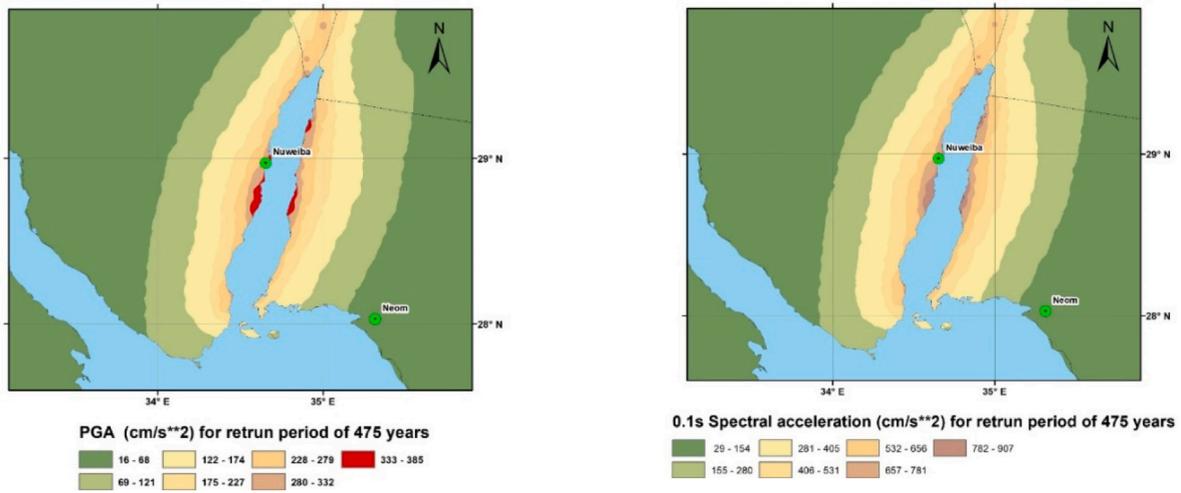
## 8. Results and discussions

The PSHA is the likelihood of ground motion exceeding a specific level at a given location within a designated return period (Bommer and Pinho, 2006). In this study, within the framework of the logic tree approach, ground motion acceleration was computed in the vicinity of the Gulf of Aqaba under bedrock conditions. The outcomes created seismic hazard curves for each point in the predefined grid of  $0.1^\circ \times 0.1^\circ$  that covers the northwestern part of Saudi Arabia and the Sinai Peninsula. Uniform hazard spectra (UHS) and seismic hazard maps were then developed using these curves. For Neom City, UHS was generated for return periods of 475, 975, and 2475 years, corresponding to probabilities of ground motion exceeding 10 %, 5 %, and 2 % over 50 years, as depicted in Fig. 6. The results indicated that the expected PGA values at Neom City are approximately  $55 \text{ cm/s}^2$ ,  $71 \text{ cm/s}^2$ , and  $95 \text{ cm/s}^2$  for return periods of 475, 975, and 2475 years, respectively. The highest values are observed at a spectral period of 0.1 s, with values of  $124 \text{ cm/s}^2$ ,  $161 \text{ cm/s}^2$ , and  $234 \text{ cm/s}^2$  for the same order of return periods. Fig. 7a-c illustrate the spatial distribution of expected PGA originating from the Gulf of Aqaba earthquake source zone. The highest value is around  $385 \text{ cm/s}^2$  for a return period of 475 years, reaching about  $577 \text{ cm/s}^2$ .

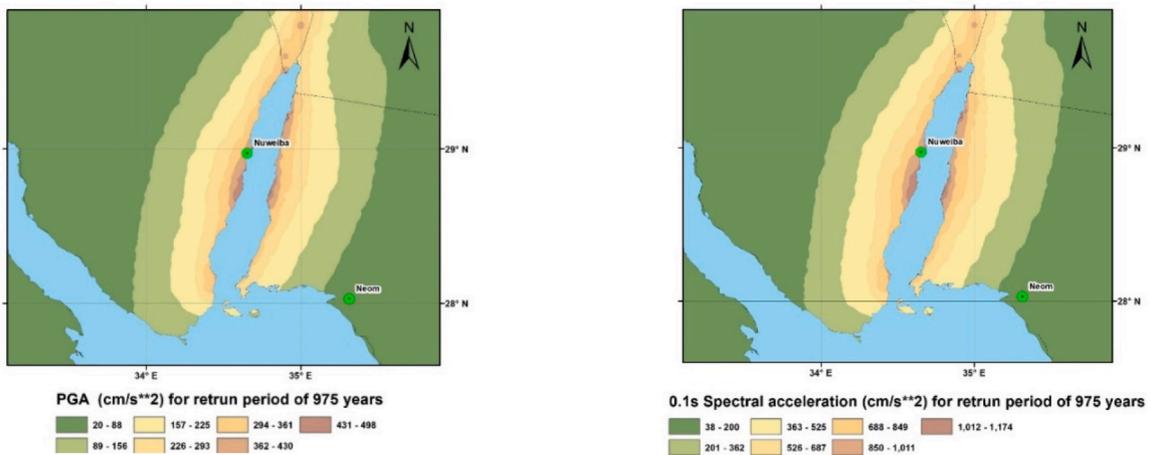
Comparisons with previous studies in the area, such as El-Eraki et al. (2015), Mostafa et al. (2018), and Abdelwahed et al. (2020), show that the estimates of PGAs are comparable. In contrast, the results of Al-Arifi et al. (2013) and Gorshkov et al. (2019) reported relatively lower PGA values, ranging from 100 to  $325 \text{ cm/s}^2$ , for a return period of 475 years. Additionally, Mostafa et al. (2018) observed PGAs at 0.2 s and 0.3 s, while in this study, the PGAs are in the range of 0.1–0.2 s, respectively. These results suggest that the Gulf of Aqaba plays a dominant role in controlling seismic hazards in the region, particularly in comparison to studies that considered multiple earthquake source zones.

## 9. Conclusions

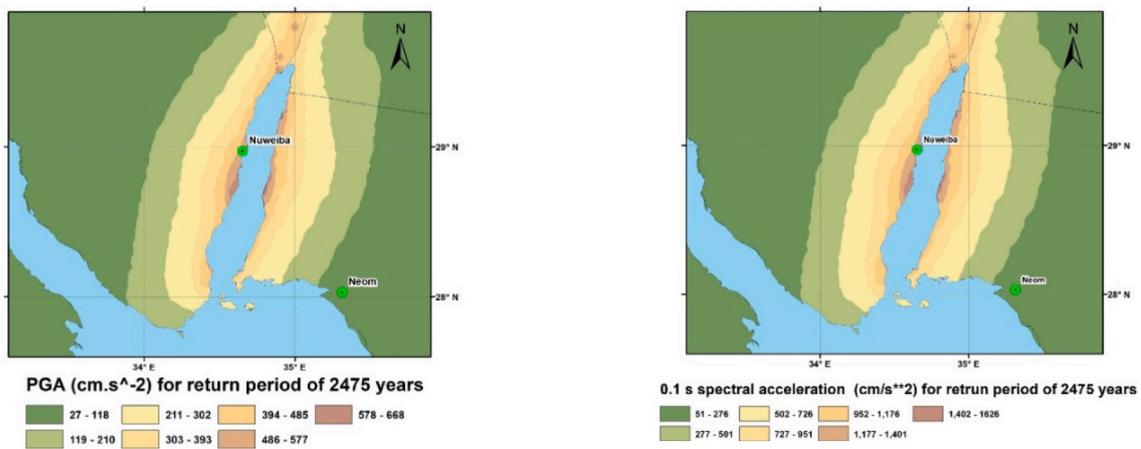
This study explores the seismic characteristics of the Gulf of Aqaba seismogenic zone and assesses the potential seismic hazard that could affect the northwestern region of Saudi Arabia. Using an updated



(a)



(b)



(c)

Fig. 7. Mean peak ground acceleration ( $\text{cm/s}^2$ ) on the rock sites, (a) with a 10 % probability of exceedance in 50 years (475-year return period), (b) with a 5 % probability of exceedance in 50 years (975 year return period), and (c) with 2 % probability of being exceeded in 50 years (2475 years return period).

earthquake catalogue developed in this study, employing a unified moment-magnitude scale, the maximum likelihood method was applied to estimate critical seismic activity parameters representing the seismogenic zone of the Gulf of Aqaba. These parameters include the maximum expected magnitude ( $M_{max}$ ), seismic hazard parameters that describe the magnitude-frequency relationship ( $\beta$ -value), and the mean seismic activity rate ( $\lambda$ ). The seismicity model proposed in this study accounts for three seismogenic zones corresponding to the Aqaba fault, the Aragonese fault, and the Arnona fault.

The current findings reveal that the maximum PGA values reach approximately  $385 \text{ cm/s}^2$  for an event expected every 475 years with a 10 % chance of occurrence,  $498 \text{ cm/s}^2$  for an event anticipated every 975 years with a 5 % likelihood, and  $688 \text{ cm/s}^2$  for an event projected every 2475 years with a 2 % chance. The response spectrum demonstrates that PGA energies exhibit high-frequency characteristics in the range of 5.0 Hz to 10.0 Hz, which raises concerns about soil response at sites near the Gulf. It is essential to avoid resonant vibrations that may coincide with the natural periods of structures.

The PSHA results indicate that the Gulf of Aqaba is a significant seismic hazard zone, particularly the Aragonese sub-seismogenic zone, where an earthquake with a moment magnitude of 7.7 is considered a realistic possibility, implying that decision-makers and policymakers in the region should consider seismic risk factors when making land use and building code decisions, especially in mega-project cities.

#### Credit authorship contribution statement

AA and ME participated in the conceptualization, analysis, calculation of seismic hazards, and writing. HB participated in the compilation of earthquake catalogue. HA participated in funding acquisition, data curation, and writing. SQ participated validation and methodology. NA participated in review and improve languistic writing of the article.

#### Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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#### Appendix A. Supplementary data

Supplementary data to this article can be found online at <https://doi.org/10.1016/j.jksus.2024.103114>.

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